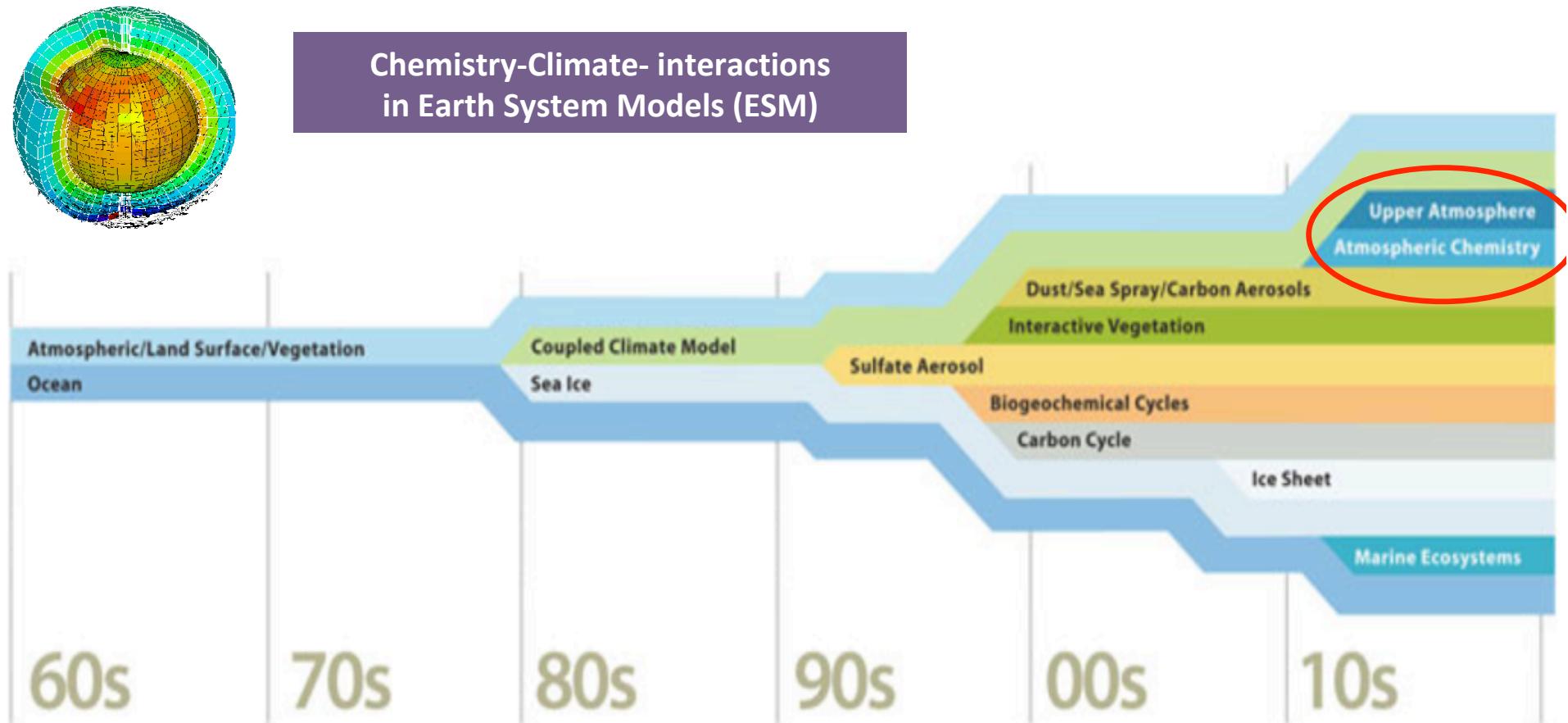


## Axe 3 : Simulation of the Cenozoic atmospheric chemistry and its interactions in the Earth System

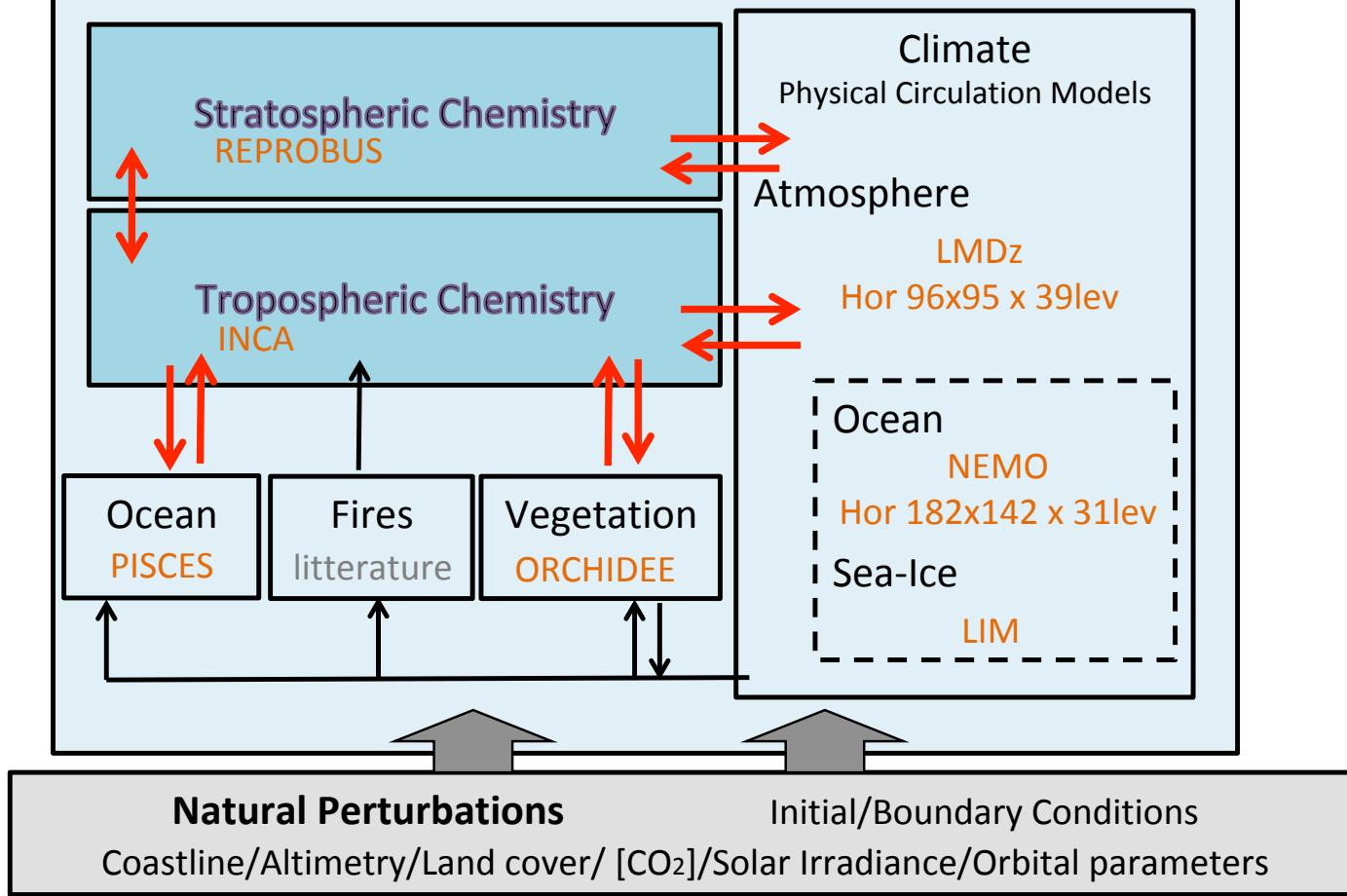


Schematic of components. Evolution of the parts of the Earth system treated in climate models over time. (Source UCAR)

⇒ Implementation of atmospheric chemistry in many of them

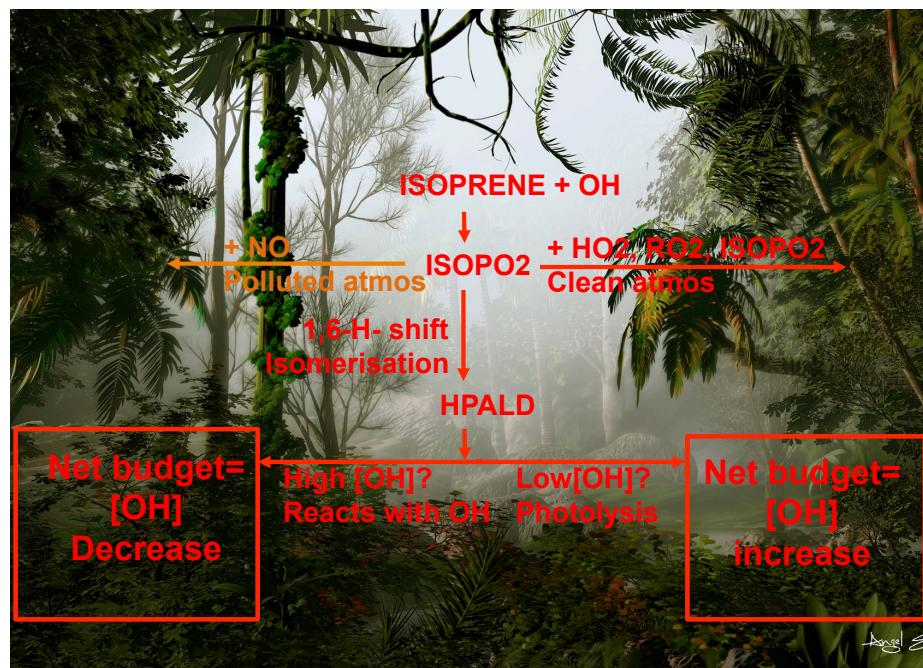
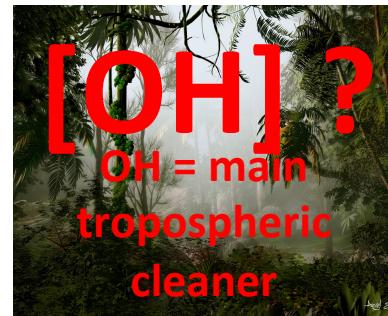
⇒ BUT NOT CONSIDERED IN PAST CLIMATE STUDIES

## IPSL-CM5-A2 Earth System Model



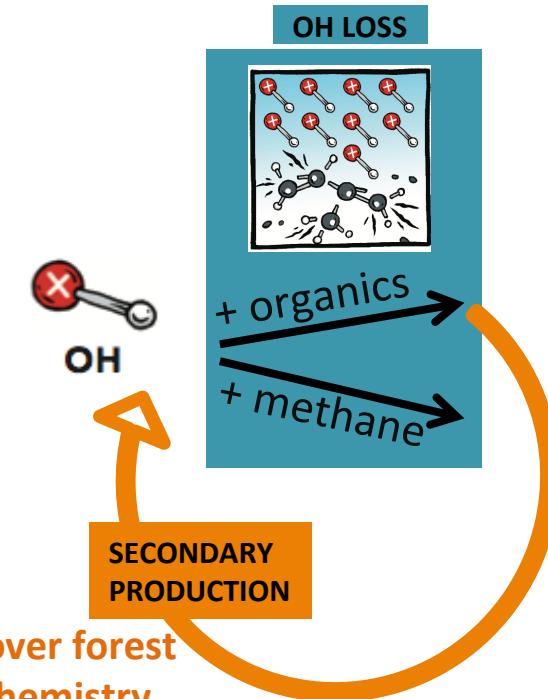
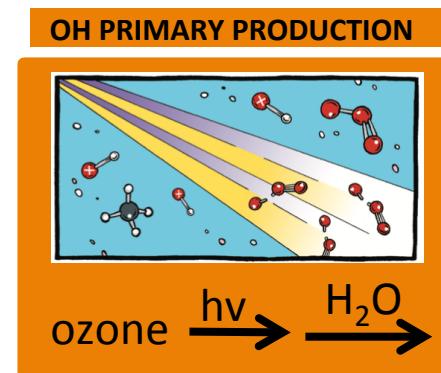
- Couplage entre chimie (strato et tropo) Xuezhou LU et Anne COZIC
- Mise en place d'une config avec ocean et autres composantes du système Terre au sein de IPSL-CM5-A2, Xuezhou LU et Anne COZIC
- Emissions océaniques dynamiques Ludivine CONTE, Laurent BOPP, SSz
- Chimie dédiée aux atmosphères faiblement polluées Cyril KARAM, SSz

## Recent progress in natural photooxidative atmospheric chemistry



TRAVAIL DE Cyril KARAM (doctorant)

## Oxidizing Capacity (=self-cleaning) ...

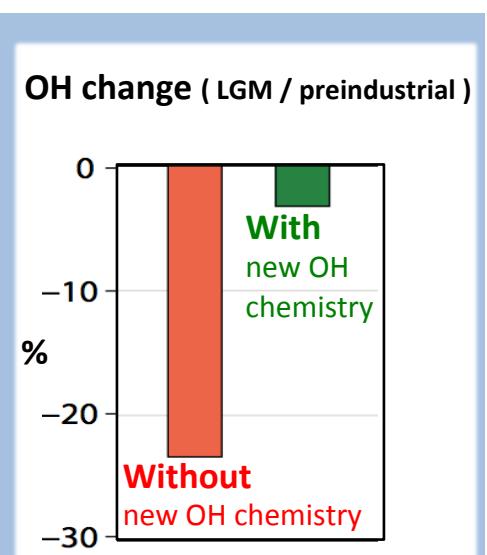


### Past COLD climates

#### Last Glacial Maximum

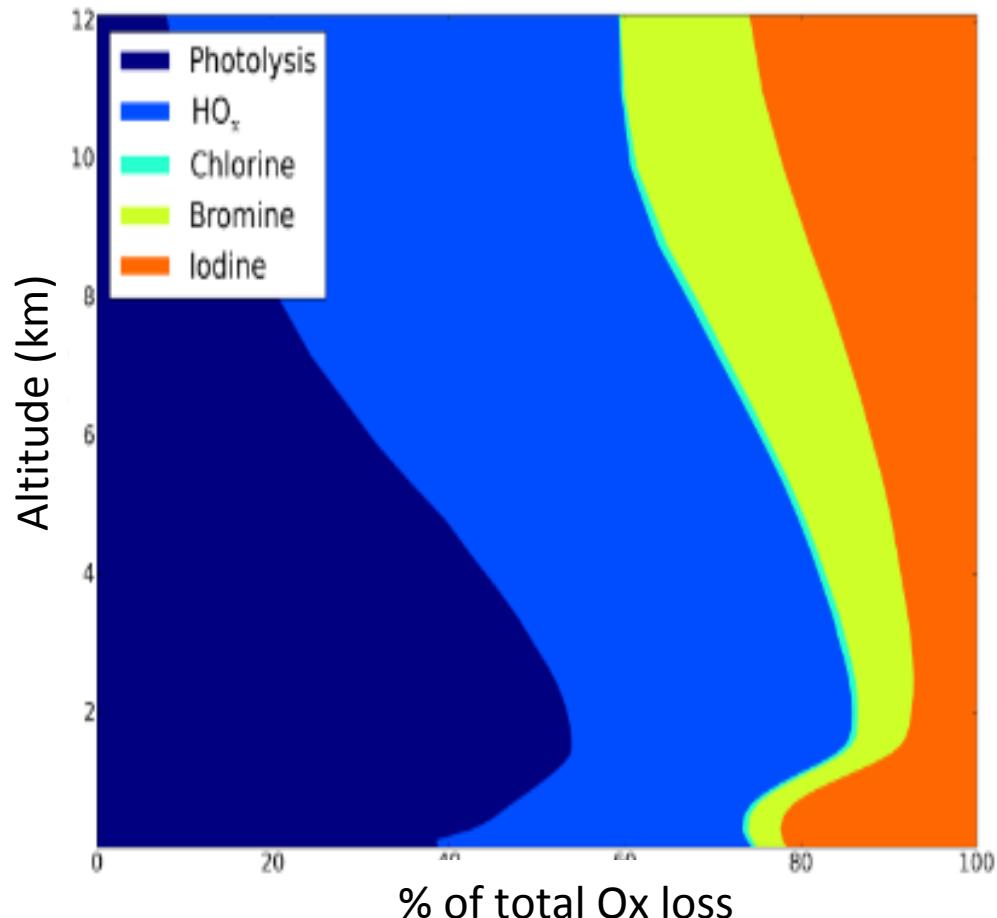
New OH chemistry dramatically alters the sensitivity of oxidizing capacity to global changes

Achakulwisut et al. 2015



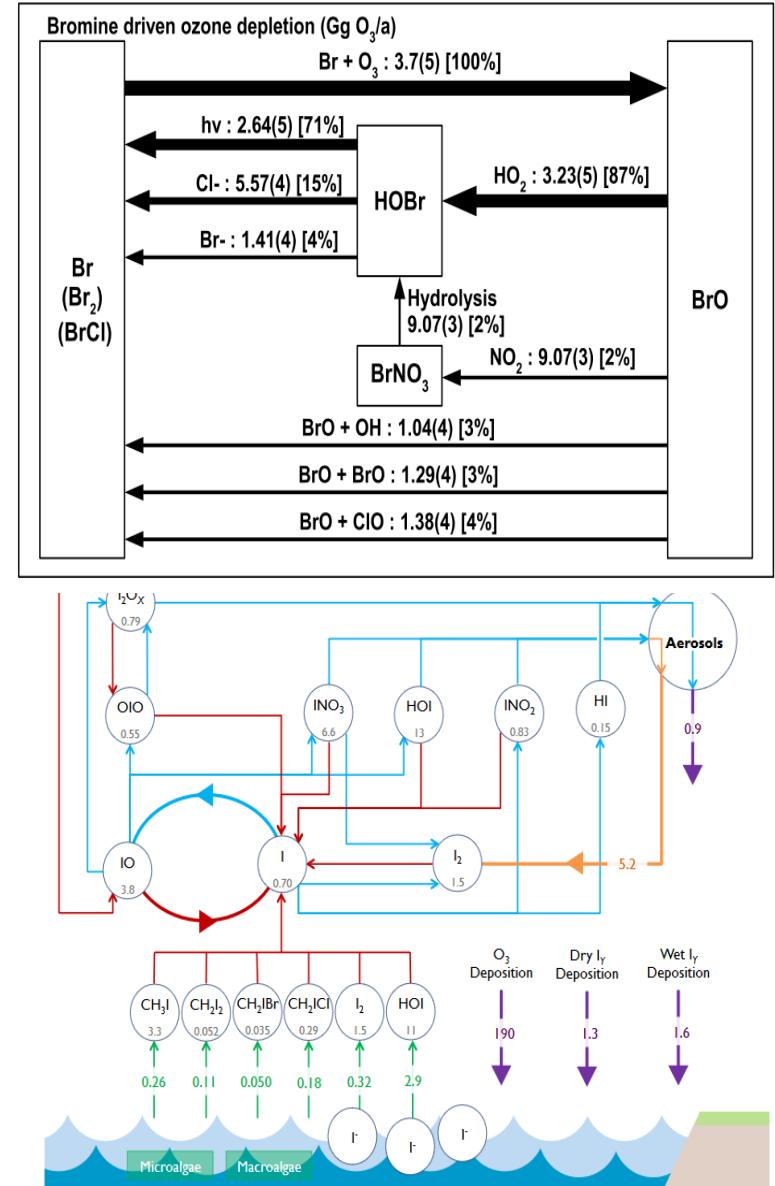
## Recent progress in natural photooxidative atmospheric chemistry

→ Importance of marine halogenated compounds in the troposphere



Global Annual-average tropospheric vertical odd oxygen loss (Ox) through different reaction routes (Photolysis,  $\text{HO}_x$ ,  $\text{I}_\text{ox}$ ,  $\text{BrO}_x$  and  $\text{ClO}_x$ )  
from Sherwen et al. 2016

Implementation de la chimie de l'Iode et du Brome  
TRAVAIL DE Cyril KARAM (doctorant)



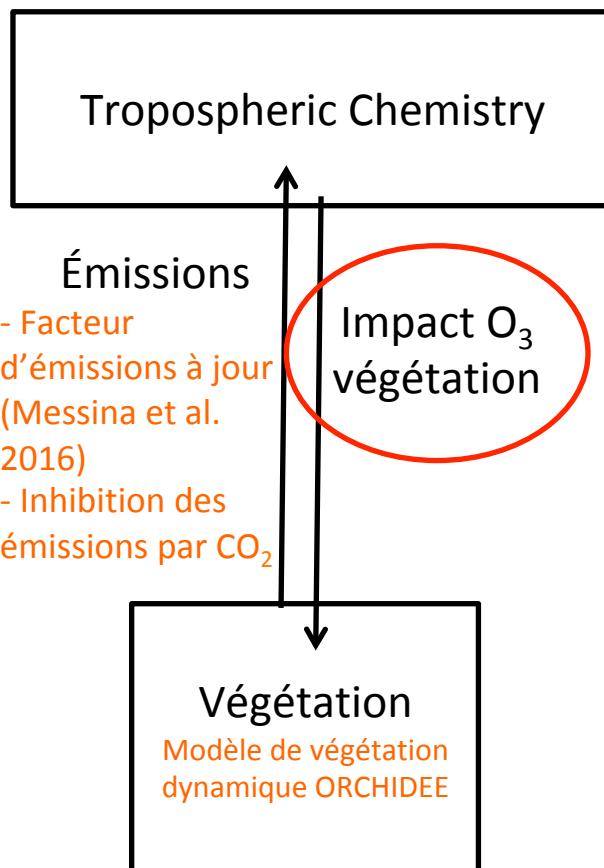
+ Other progress in chemistry of oxidants (Criegee, peroxy radicals, etc.)

# Perspectives: Interactions biosphère /chimie atmosphérique



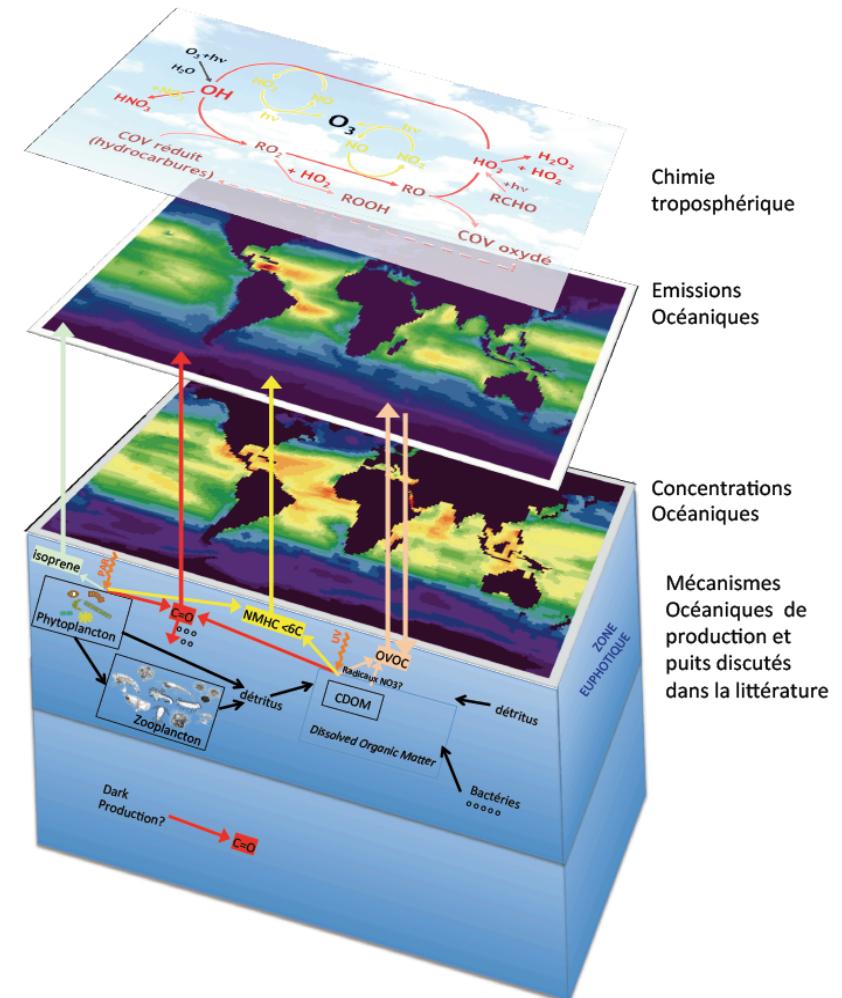
avec biosphère terrestre

J. Lathiére



avec biosphère marine

L Conte/ L Bopp



## Methods

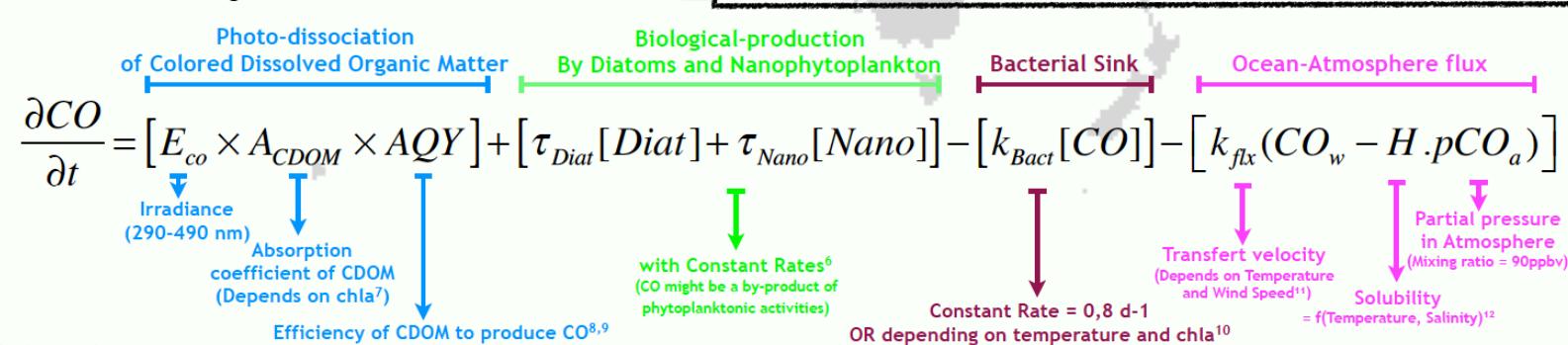
### NEMO-PISCES model :

- A global circulation model (NEMO) and a Biogeochemical model (PISCES)<sup>17</sup>
- 2 phytoplankton groups / 3 non-living organic matter pools / 5 nutrients
- Run for 3000 years to reach equilibrium using a mean climatology
- Resolution :  $2^\circ \times 2^\circ \cos(\text{lat})$  with 31 vertical levels

## Dynamical oceanic emissions of tropospheric chemistry reactants

TRAVAIL DE Ludivine CONTE (doctorante)

### Processes Affecting oceanic CO concentration :



- We used in situ measurements of surface CO concentration to evaluate simulated concentrations

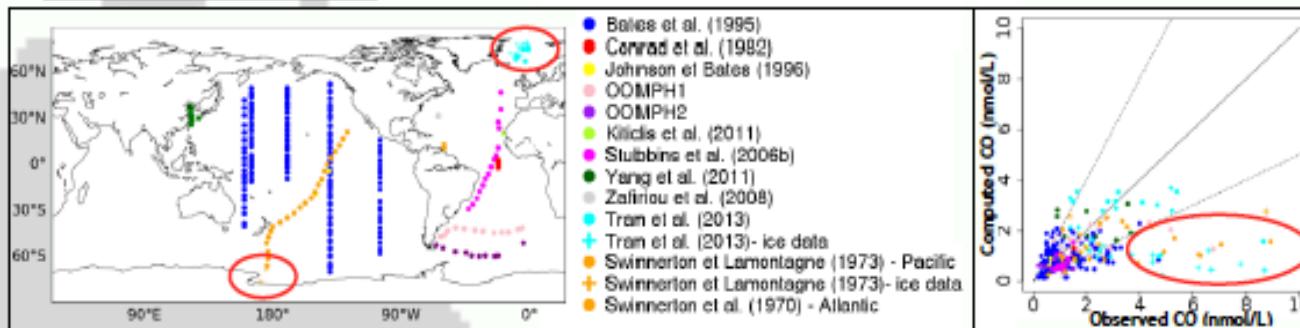


Figure 2 : Left : Position of in situ measurements of surface CO concentrations, used to evaluate computed concentrations. Right : scatter plot of computed versus observed surface CO concentrations for a run with a mean bacterial consumption rate of 0,8d<sup>-1</sup> (RMSE = 1,38). Colors refer to the origin of the data. Red circles show polar data.

- High in situ data (up to 9 nmol/L) in polar regions (red circles), are not accurately represented
- Those polar data present a large spatial and temporal variability and suggest specific CO production involved in ice-covered regions

## Dynamical oceanic emissions of tropospheric chemistry reactants

# How Emissions respond to Climate Change ?

- We performed two long runs (from 1850 to 2100) :
  - => one with a constant pre-industrial atmospheric CO<sub>2</sub> concentration (285 ppm)
  - => one with an increasing CO<sub>2</sub> concentration corresponding to RCP 8.5 (up to 936 ppm )
- The global CO flux increases by 15,6% under an increasing atmospheric CO<sub>2</sub> (Figure 5)
- The global budget of biological sources and sink show little change

### Attribution of the rise of the global CO flux to a combination of physical processes :

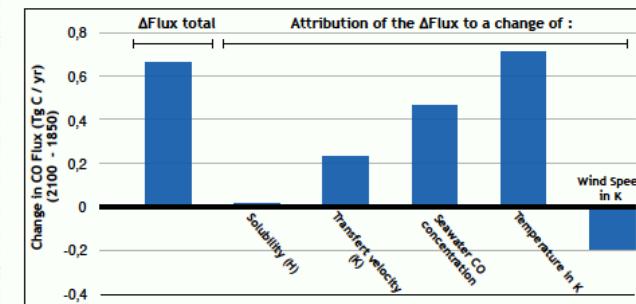


Figure 6 : Change in the ocean-atmosphere CO flux (TgC/yr) : calculated by PISCES or recalculated offline : if only H, K, or CO are changed in 2100 compared with 1850. A positive change means an increased flux in 2100 under climate change.

=> Solubility of CO (*H*) decreases due to a global rise of sea surface temperature

=> Transfert velocity (*k<sub>flux</sub>*), depending on temperature and wind speed, increases

=> CO concentration in surface seawater tends to increase slightly :  
Temperature ↗ => stratification ↗ => mixing ↘

=> Among those processes, the change of the gaz transfert velocity *k<sub>flux</sub>* due to the surface temperature increase dominates to increase the CO flux towards the atmosphere (Figure 6)

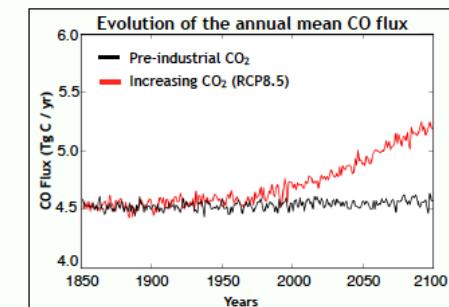
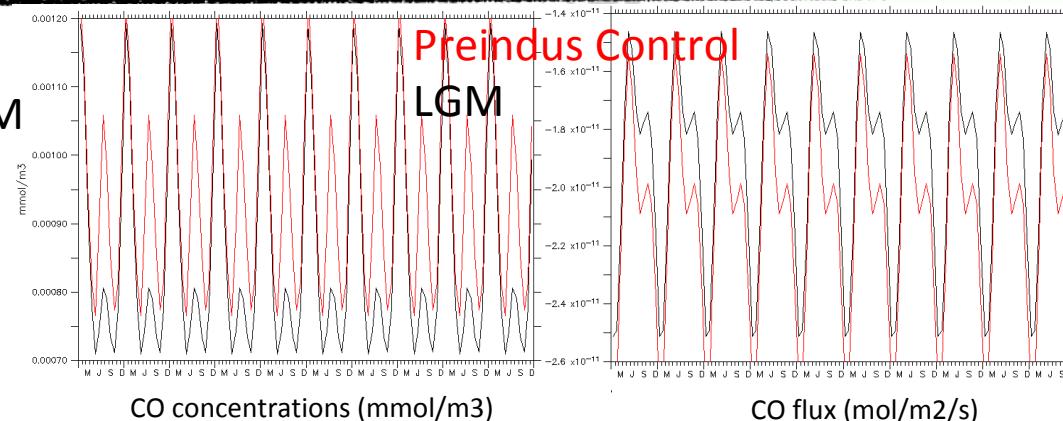


Figure 5 : Temporal evolution of the mean annual CO flux (TgC/yr). Red: with climate change under RCP8.5 scenario. Black : with constant pre-industrial CO<sub>2</sub> concentration. Positive flux is toward the atmosphere.

Tests en conditions LGM

TRAVAIL DE  
Ludivine CONTE  
(doctorante)



Concentrations plus faibles mais export vers atmosphère plus grand

lié à augmentation des vents LGM(?)

	Réunions	dec	jan	fev	mars	avr	mai	juin	juil	aout	sept	oct	nov	dec	jan	fev	mars	avr	mai
<b>Axis 2 - Development for past pristine atmospheres</b>																			
<b>WP3 Chemistry Model</b>																			
Full tropo and strato model to test on present day conditions																			
Present day climatology to be compared with observations																			
Preindustrial climatology to be compared with observations and to multimodel experiment (few data)																			
Evaluation of the model performance for present-day and preindustrial conditions																			
<b>WP4 New natural emissions</b>																			
Evaluation of Biogenic emission range (from ORCHIDEE) for each scenario																			
Evaluation of Oceanic emission range (from PISCES) for each scenario																			
Evaluation of Wildfires emission range (from litterature and collaboration) for each scenario																			
Sensitivity studies to Natural emissions																			

En retard car stratégie un peu modifiée (tests sur run eocene avant tout)

(+ thèses C Karam et Ludivine Conte) mais pas de problèmes majeurs

Devlpts techniques ont bien avancé

A VENIR 12 prochains mois : evaluation du modèle de chimie et sensibilité des émissions naturelles. devrait etre OK

	Réunions	dec	jan	fev	mars	avr	mai	juin	juil	aout	sept	oct	nov	dec	jan	fev	mars	avr
<b>Axis 3 - Cenozoic atmospheric chemistry simulations</b>																		
<b>WP5 Paleo chemistry-Climate simulations</b>																		
Last Glacial Maximum Simulation																		
Eemian Simulation																		
Mid-Pliocene Simulation																		
Optimum Miocene Simulation																		
Paleocene-Eocene Thermal Maximum Simulation																		
Analysis of the simulations, realism of the results, comparison with previous study and ISOTOPE																		
Climatologies of 3D distribution of reactive compounds and corresponding surface UV radiation																		
<b>WP6 Earth System Feedbacks</b>																		
Simulation of the climate feedback due to composition change (ESM forced by WP5 concentrations) for each of the 6 past conditions																		
Quantification of chemistry effect on climate																		
IPSL-CM5 model with interactions between: Climate and 3D atmospheric N2O, CH4, O3/Chemistry and terrestrial biosphere/Atm Chemistry and marine biogeochem																		
Simulation with the IPSL-CM5 model first with all the couplings for hot climate conditions (100 yrs + 1base line 100yr)																		
Quantification of the feedbacks																		

WP5 : priorité a été mise sur tester config Eocene, pas de probleme non résolus

Les 4 autres configs passés vont etre testés en LMDz-INCA/ LMDz-reprobus ET LMDzINCA\_REPR d'ici a la fin de l'été grace aux 300000h supp obtenues sur curie

WP6 : on y pensera plutot en 2019