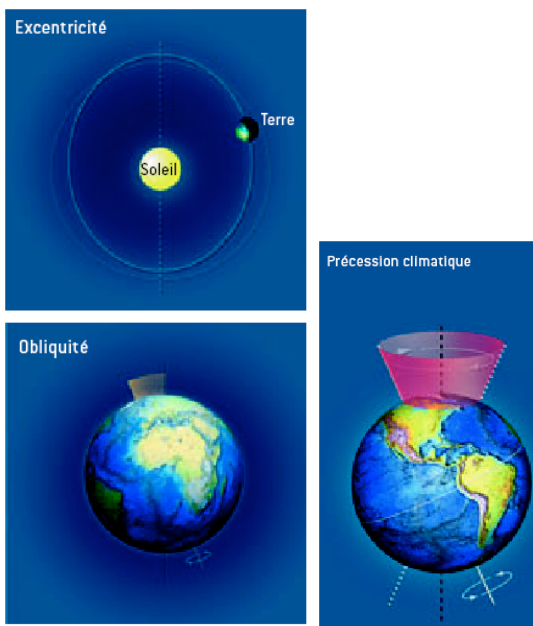


## Glacial-Interglacial cycles : Insights from a conceptual model

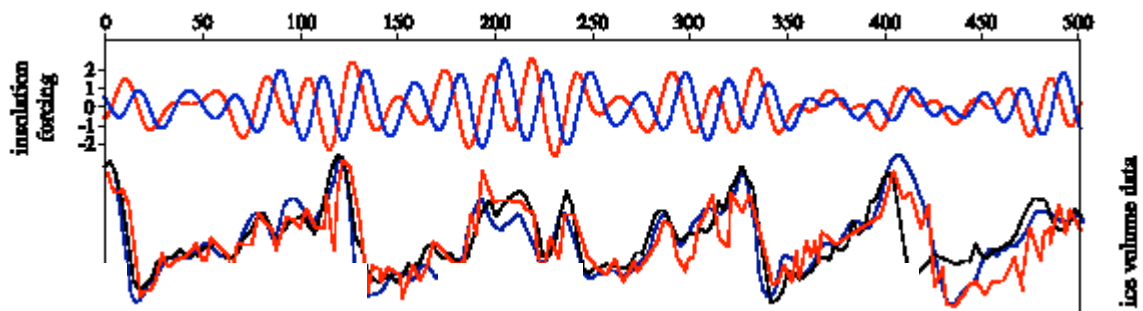
### Introduction :

**Milankovitch theory** : Northern Summer Insolation, driven by astronomical parameters (precession, obliquity and excentricity), is the key driver of the interglacial-glacial cycles (astronomical theory of quaternary climate).



Periodicity of those processes  
Precession : 26000 yrs  
Obliquity : 41 000 yrs  
Precession : 100 000 and 400 000 yrs

### Observations :

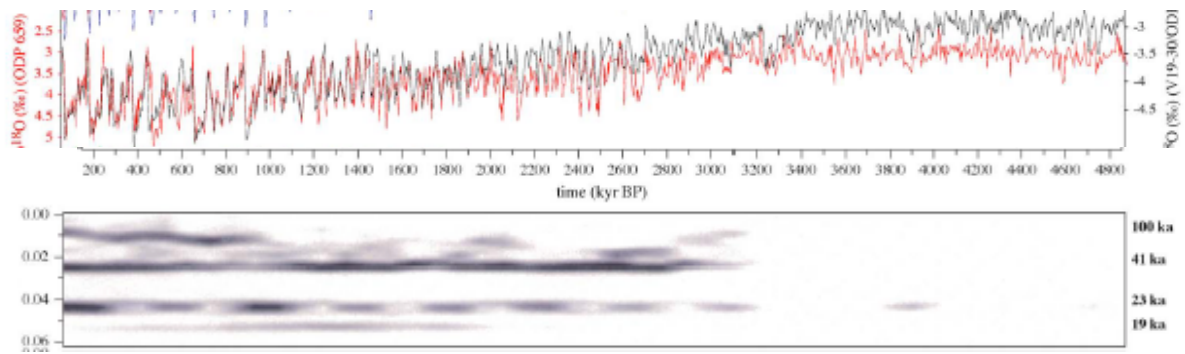


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### Open questions :

100 000 yr cycle ?

Relation between insolation and ice volume not so clear ? Where are the non-linearities ?  
 Role and explanation of atmospheric CO<sub>2</sub> variations ?  
 Changes in the main frequency over the last millions years ?



Another theory to explain glacial-interglacial cycles of CO<sub>2</sub> and climate



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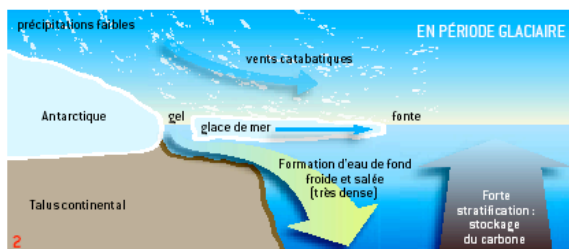
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## The Antarctic ice sheet and the triggering of deglaciations

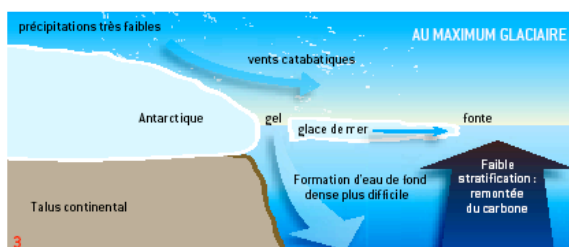
Didier Paillard<sup>a,\*</sup>, Frédéric Parrenin<sup>b,1</sup>



Interglacial :  
 Formation of fresh and cool water.  
 Little stratification.  
 CO<sub>2</sub> escapes.



Glacial :  
 Less precipitation, more sea ice  
 Formation of very salty and cool water.  
 Strong stratification.  
 CO<sub>2</sub> stays in the deep ocean.



Glacial Maximum:  
 Antarctic Ice cap everywhere over the shelf  
 Formation of deep water is difficult.  
 Little stratification.  
 CO<sub>2</sub> escapes.

## How do we test this theory ?

### By a conceptual Model and its mathematical formulation

A mathematical expression of this conceptual model follows. We need three variables: global ice volume  $V$  forced both by Northern Hemisphere summer insolation and atmospheric  $p\text{CO}_2$ ; extent of Antarctic ice sheet  $A$  forced by sea level changes (i.e., by  $V$ ); and atmospheric  $\text{CO}_2$   $C$  linked primarily to deep-water state «glacial» or «interglacial». The oceanic switch is forced by the «salty bottom waters formation efficiency» parameter:

$$F = aV - bA - cI_{60} + d$$

which increases with  $V$  and decreases with  $A$ .  $a$ ,  $b$ ,  $c$ , and  $d$  are constant coefficients. Indeed,  $F$  should increase when global climate cools (through  $V$ ) and decrease when continental shelf areas are reduced (through  $A$ ).  $I_{60}$  is the daily insolation ( $60^\circ\text{S}$ , 21st February) inasmuch as a reduced sea-ice extent during late Austral summer could affect the Southern Ocean heat budget, and consequently, warm the regional climate, thus affecting  $F$ . Coefficient  $c$  is typically very small. When  $F$  is negative, the ocean is in «interglacial» mode and reciprocally. The model equations are:

$$dV/dt = (V_R - V)/\tau_V$$

$$dA/dt = (V - A)/\tau_A$$

$$dC/dt = (C_R - C)/\tau_C$$

with  $V_R = -x C - y I_{65} + z$  (i.e., ice volume is driven by insolation and  $\text{CO}_2$ );  $C_R = \alpha I_{65} - \beta V + \gamma H(-F) + \delta$  (i.e.,  $\text{CO}_2$  is driven by some precessional forcing, here  $I_{65}$ , by global climate  $V$  and by deep ocean stratification);  $H$  is the Heaviside function ( $H=1$  if  $F < 0$ ;  $H=0$  otherwise);  $I_{65}$ =insolation  $65^\circ\text{N}$ , 21st June;  $\tau_V$ ,  $\tau_A$ ,

and  $\tau_C$  are time constants;  $x$ ,  $y$ ,  $z$ ,  $\alpha$ ,  $\beta$ ,  $\gamma$ , and  $\delta$  are constants. Contribution of Antarctica to  $V$  is neglected.  $V$  is limited by a minimum zero value. This system is linear, except for the discontinuity represented by  $H(-F)$ , which reflects either a non-linearity in the carbon cycle, or more probably, a nonlinearity of the interactions between deep stratification, bottom water formation, and thermohaline

**3 questions :**

- **Is this simple conceptual model able to catch the main features of the glacial-interglacial cycles when forced by insolation ?**
- **Is this simple model able to catch the changes in frequencies over the last 5 millions years ?**
- **“What if” question : If we add a massive amount of carbon in the atmosphere, when does this simple model predict the next glaciation ?**