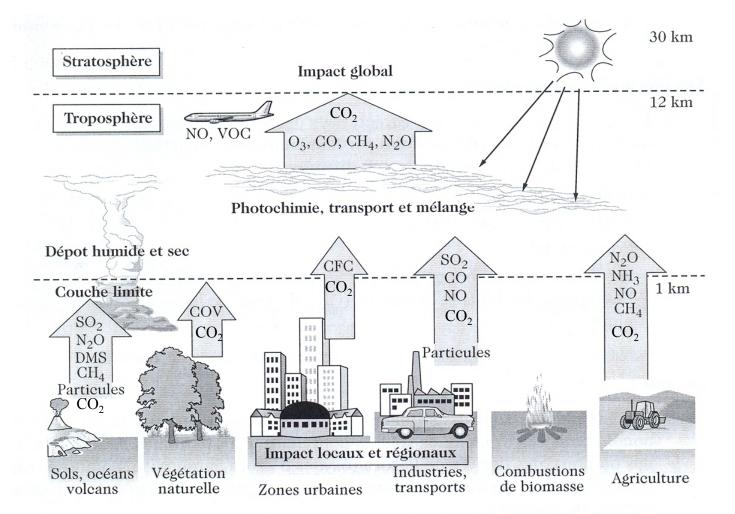
The budget of tropospheric ozone and its evolution

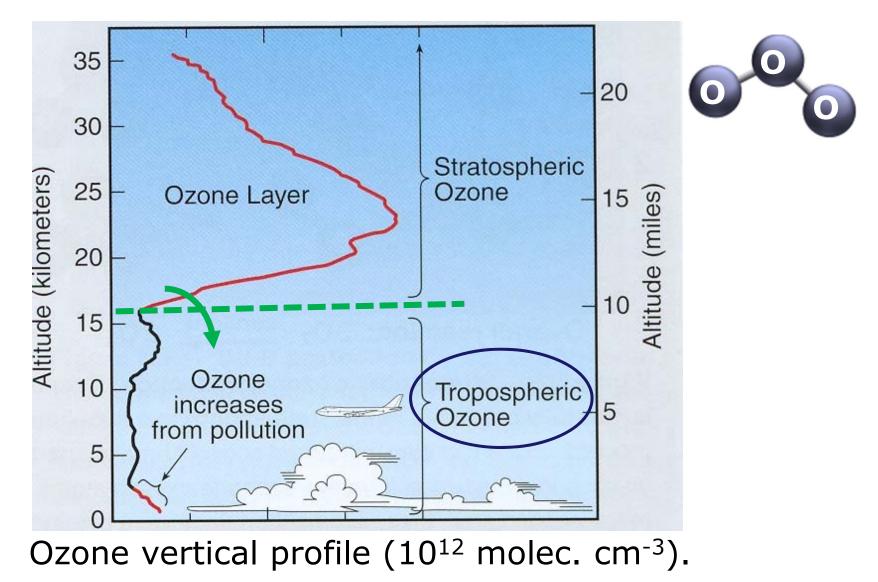
Didier Hauglustaine

Laboratoire des Sciences du Climat et de l'Environnement (LSCE), Gif-sur-Yvette Laboratoire Image Ville Environnement (LIVE), Strasbourg



Major processes controlling the atmospheric composition.

About 90% of the ozone molecules contained in the atmosphere are located in the stratosphere



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Formation and destruction of the hydroxyl radical OH

Photochemical source of OH

 $\begin{array}{rl} O_3 + h_{\mathcal{V}} \rightarrow & O_2 + O(^1D) & (R1) \\ & (\lambda < 310 \text{ nm}) \end{array}$

 $O(^{1}D) + M \rightarrow O(^{3}P) + M$ (R2)

 $O(^{1}D) + H_{2}O \rightarrow 2 OH$ (R3)

The OH concentration is controlled by the ozone distribution (O_3) , water vapour (H_2O) and solar radiation (ultraviolet).

Global average: [OH] = 1.0x10⁶ molecules cm⁻³ (10⁻¹⁴ in volume mixing ratio) OH is the major tropospheric oxidant and hence controls the residence time of most pollutants in the atmosphere.

 $CO + OH (+O_2) \rightarrow CO_2 + HO_2 (R4)$

 $CH_4 + OH (+O_2) \rightarrow CH_3O_2 + H_2O (R5)$

 $RH + OH \rightarrow R + H_2O$ (R6)

 $HCFC22 + OH \rightarrow Products (R7)$

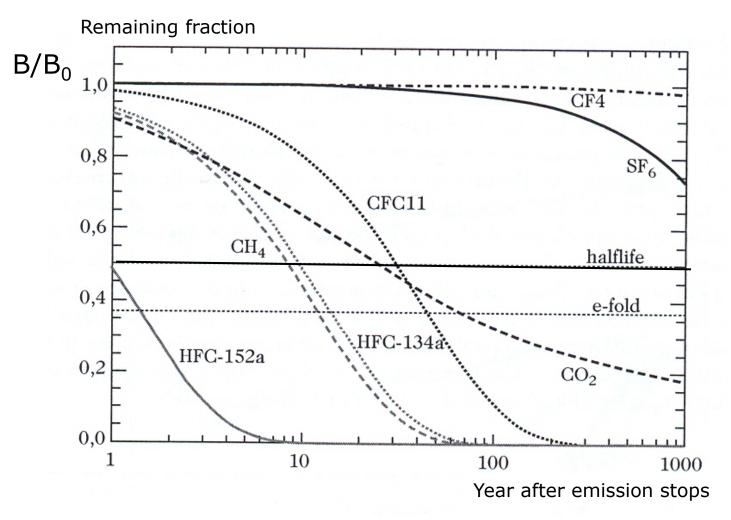
 $SO_2 + OH \rightarrow (...) \rightarrow sulfates (R8)$

Residence time of constituents in the atmosphere:

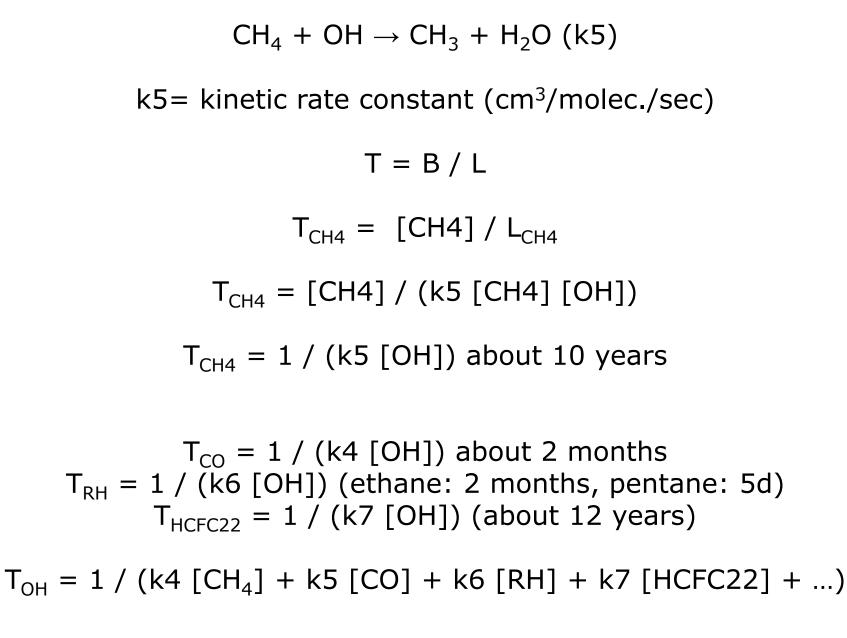
$$dB/dt = Emission - Loss = E - B/T$$

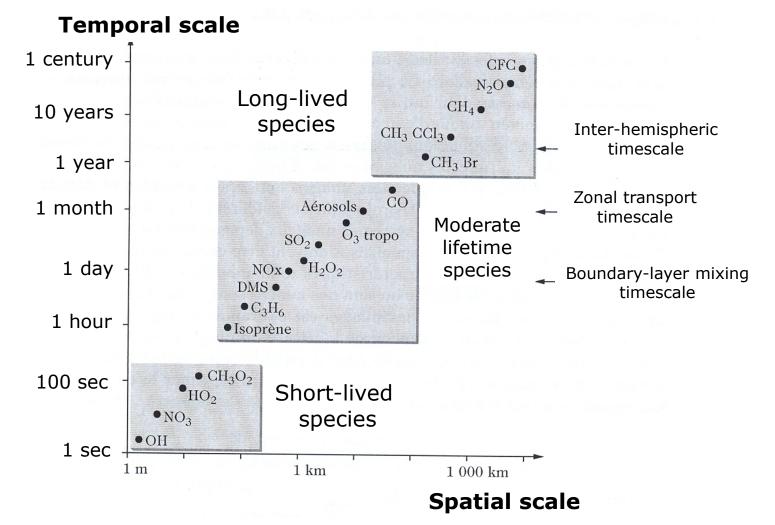
Solution when emissions go to zero :

 $B(t) = B_0 \exp(-t/T)$

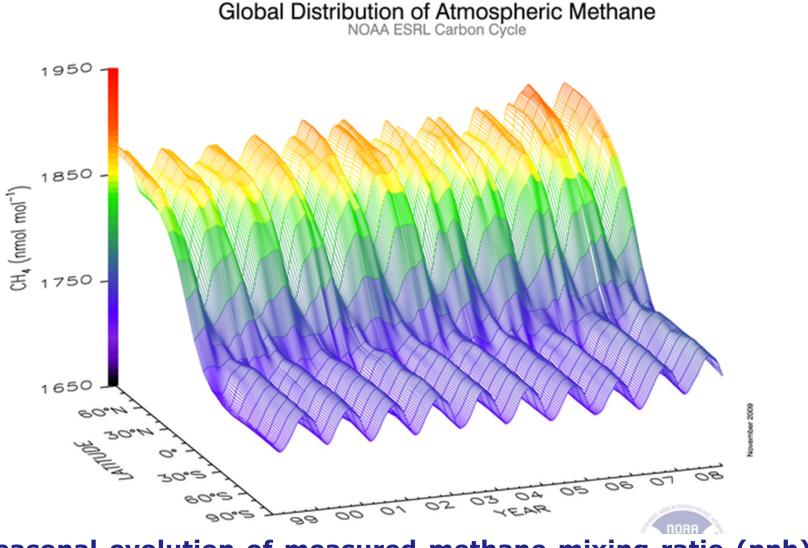


Evolution of the remaining mass of constituents in the atmosphere when their emission stops.





Correlation between the residence time of constituants and the spatial scale of their distribution.



Seasonal evolution of measured methane mixing ratio (ppb) over the 1999-2008 period.

While domestic emission controls on $NO_{x'}$ NMHC and CO combined are clearly most effective for lowering domestic $O_{3'}$ the O_3 response to anthropogenic emissions of CH₄ from distant source regions is nearly as large as that to emissions of the traditional O_3 precursors in these regions.

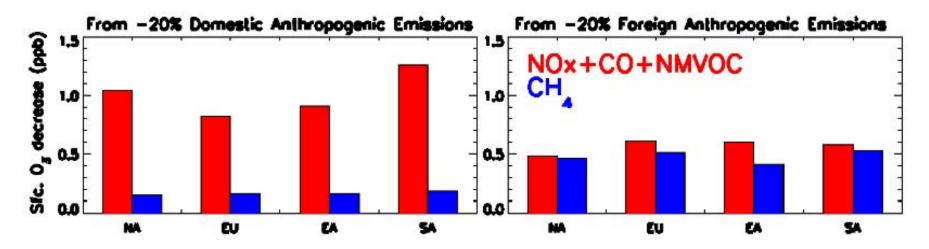
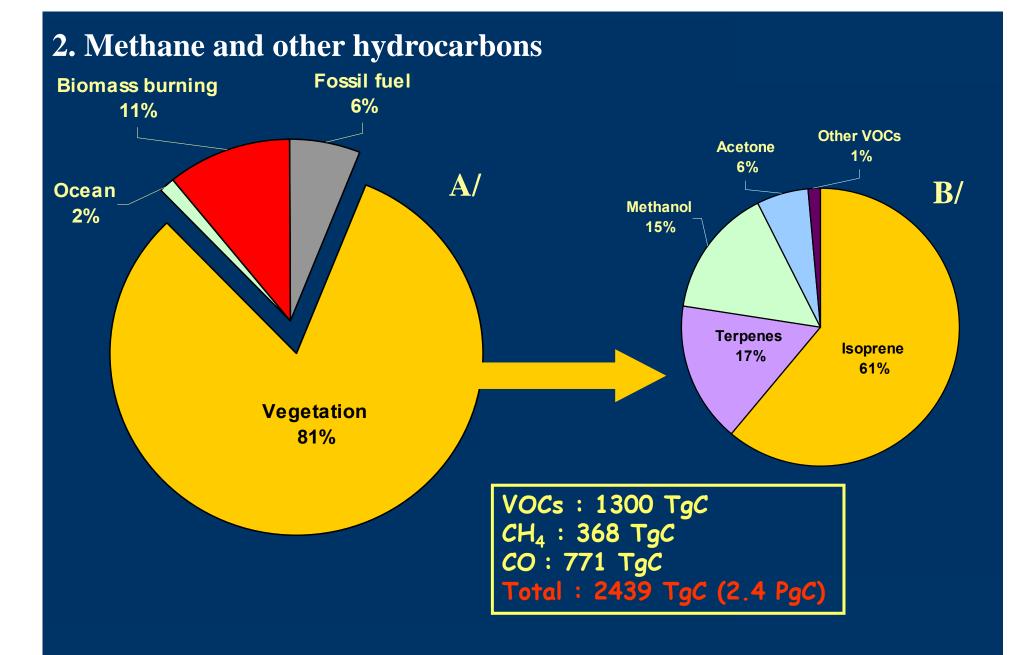
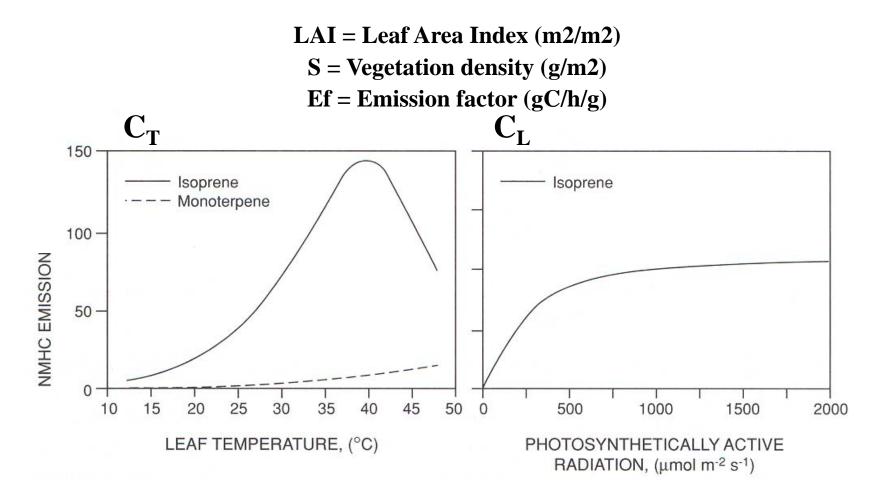


Figure 4.9. Model ensemble surface O_3 decrease (ppbv), annually and spatially averaged over the HTAP regions from 20% decreases in anthropogenic emissions of NO_x , CO and NMVOC (red) versus 20% decreases in anthropogenic CH₄ (blue). Influence of each source region on surface O_3 within the same region (termed "domestic", left panel), and the sum of the O_3 responses to emission changes within the three other source regions (termed "foreign", right panel). Adapted from Fiore et al. [2010].

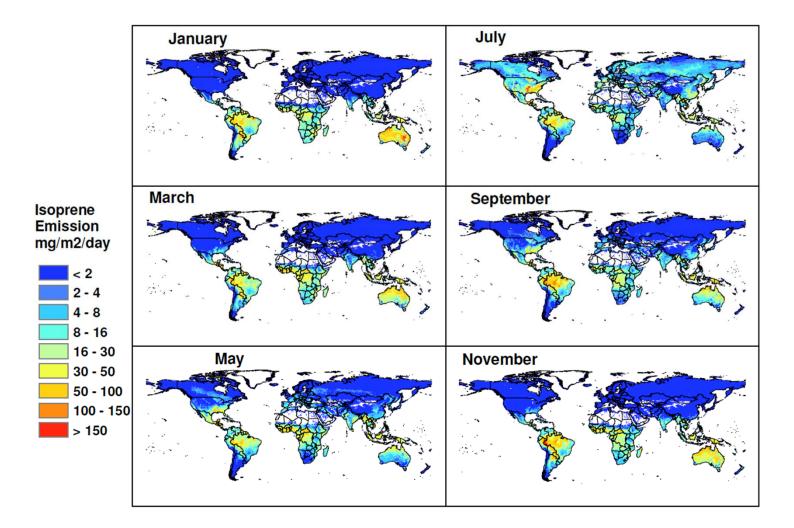


Total (A) and biospheric (B) emissions of non-methane hydrocarbons (%).

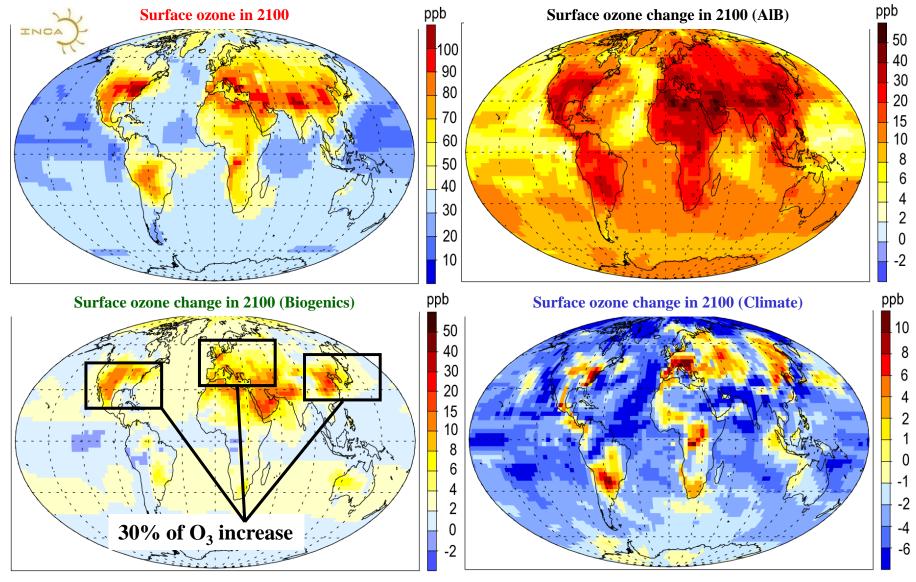
 $F(gC/m^2/h) = LAI \times S \times E_f(PFT,COV) \times C_T(T,COV) \times C_L(PAR)$



Temperature (C_T) and radiation (C_L) dependence of isoprene and monoterpenes emissions.

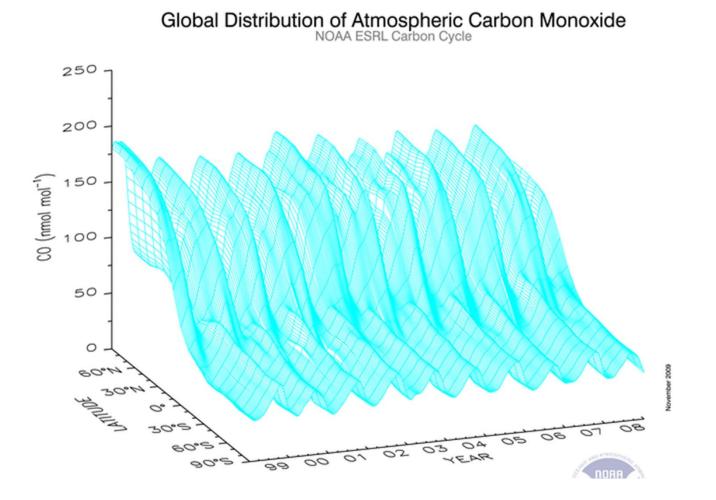


Isoprene emissions for various months simulated by the MEGAN model.



Hauglustaine et al., 2005

3. Carbon monoxide



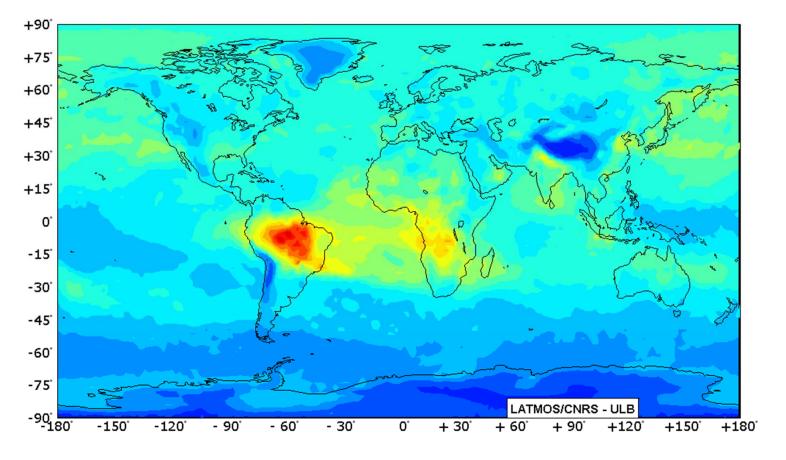
Seasonal evolution of measured CO mixing ratio (ppbv) over the 1999-2008 period.

3. Carbon monoxide

Carbon monoxide measured by IASI on board (METOP)

2008 - 2009

NOV DEC JAN FEB MAR APR MAY JUN JUL AUG SEP OCT

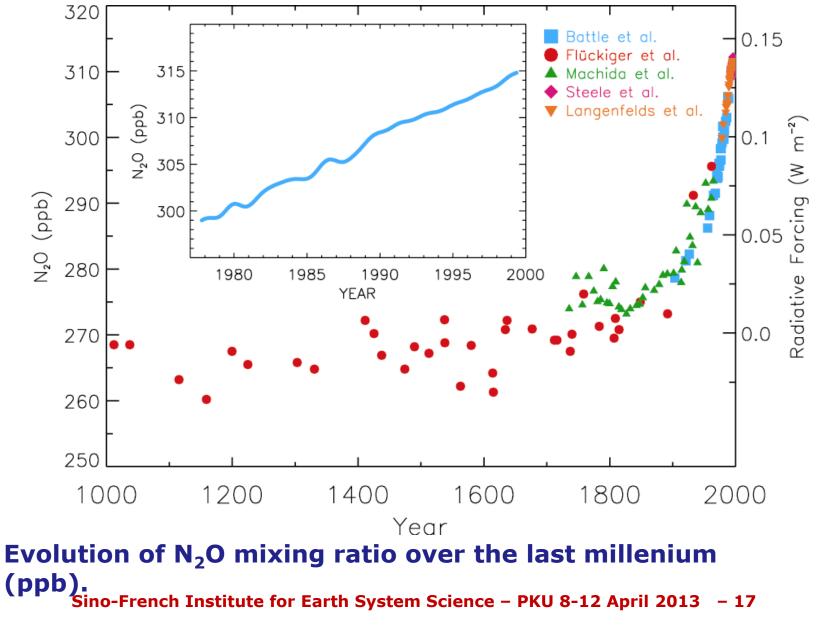




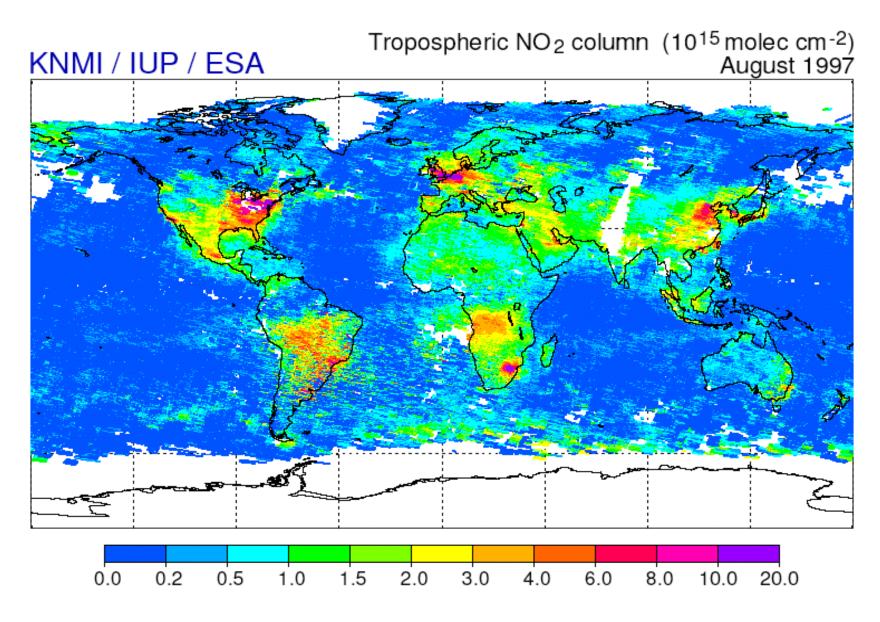
0 0.2 0.4 0.6 0.8 1 1.2 1.4 1.6 1.8 2 2.2 2.4 2.6 2.8 3 3.2 3.4 3.6 3.8 4 Total Column CO x10¹⁸ molecules/cm²

4. Nitrogen compounds

In the stratosphere: $N_2O + O(^1D) \rightarrow NO + NO$



4. Nitrogen compounds



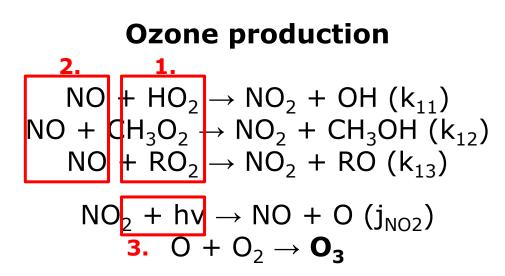
Tropospheric ozone photochemical production

Photostationnary state O_3 -NO-NO₂ (T \approx 1min)

During daytime

 $\begin{array}{l} \mathsf{NO} + \mathsf{O}_3 \rightarrow \mathsf{NO}_2 + \mathsf{O}_2 \ (k_{10}) \\ \mathsf{NO}_2 + \mathsf{hv} \rightarrow \mathsf{NO} + \mathsf{O} \ (j_{\mathsf{NO2}}) \\ \mathsf{O} + \mathsf{O}_2 \rightarrow \mathsf{O}_3 \end{array}$

 $NO_2/NO = k_{10} O_3 / j_{NO2}$

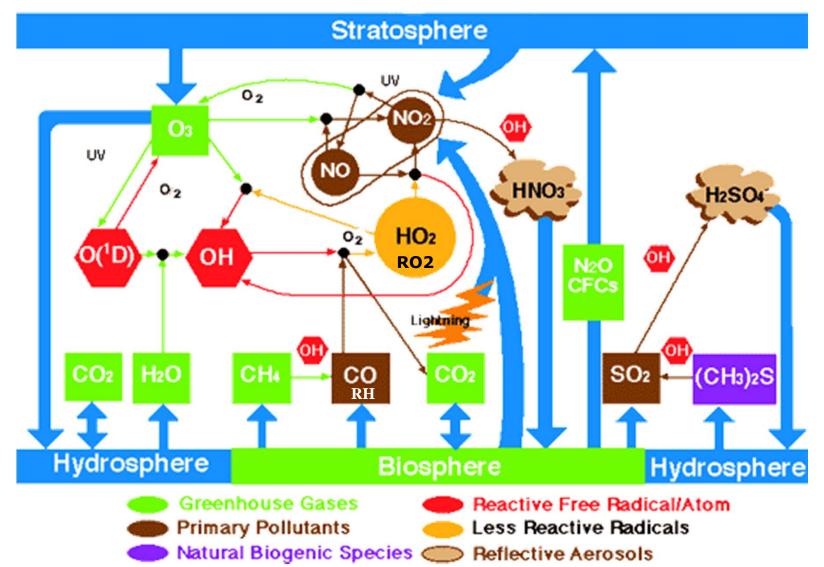


Ozone photochemical destruction

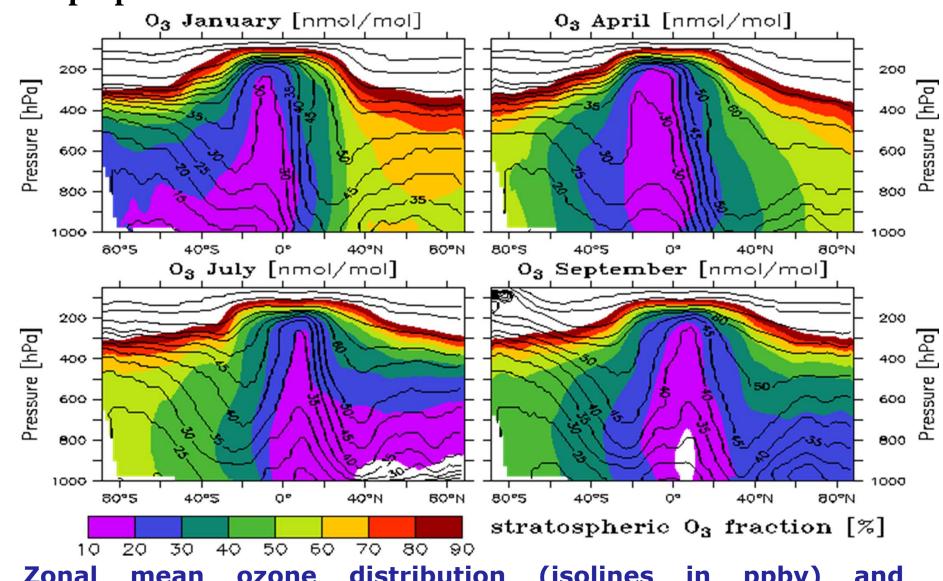
1.
$$\mathbf{O_3}$$
 + hv \rightarrow O₂ + O(¹D) (λ < 310 nm)
O(¹D) + H₂O \rightarrow 2 OH

2.
$$O_3 + HO_2 \rightarrow OH + 2 O_2$$

3.
$$\mathbf{O_3} + \mathrm{OH} \rightarrow \mathrm{HO}_2 + \mathrm{O}_2$$



Summary of major processes controlling the tropospheric ozone budget and the hydroxyl radical OH.



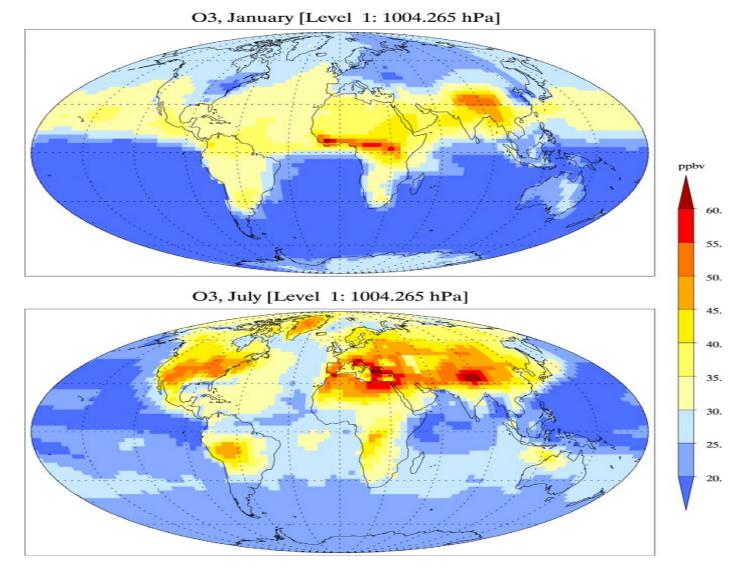
Zonal mean ozone distribution (isolines in ppbv) and stratospheric ozone fraction (%) calculated in January, April, July and September.

Tropospheric ozone budget simulated by global tropospheric models (ensemble mean).

Ozone budget terms	Tg(O ₃)/yr
Sources	
Photochemical production (P)	3948 ± 761
Stratospheric Influx	636 ± 273
Sinks	
Photochemical destruction (L)	3745 ± 554
Surface deposition	902 ± 255
Net photochemistry (P-L)	245 ± 346
Burden	307 ± 38
T = B / L	

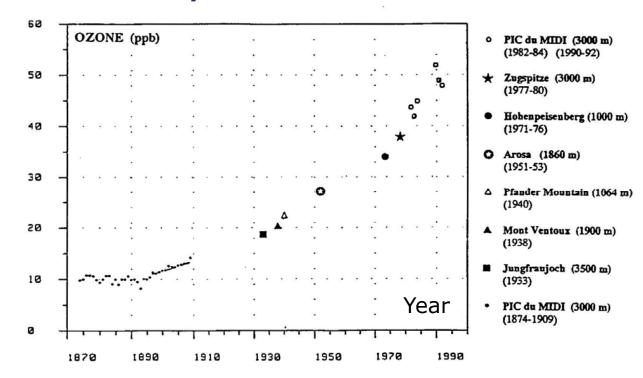
$$T = 307 / (3745+902) = 24 \text{ days}$$

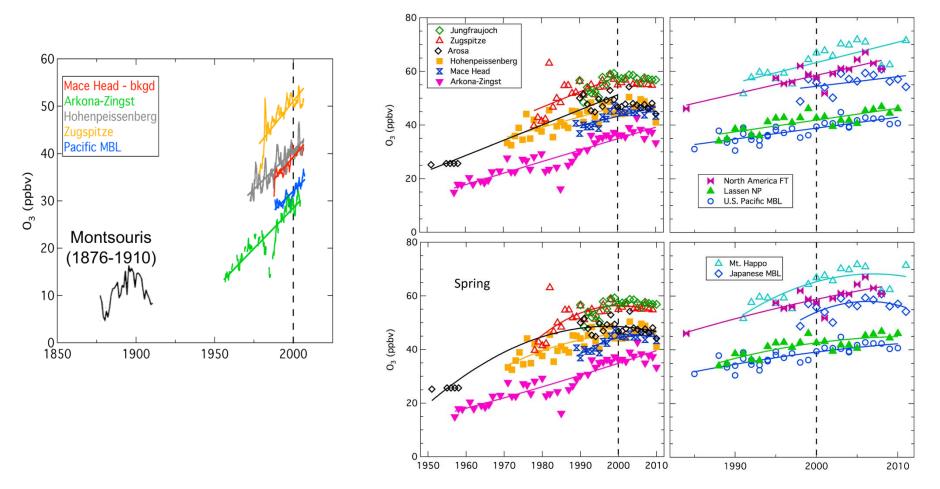
Wild et al., 2012



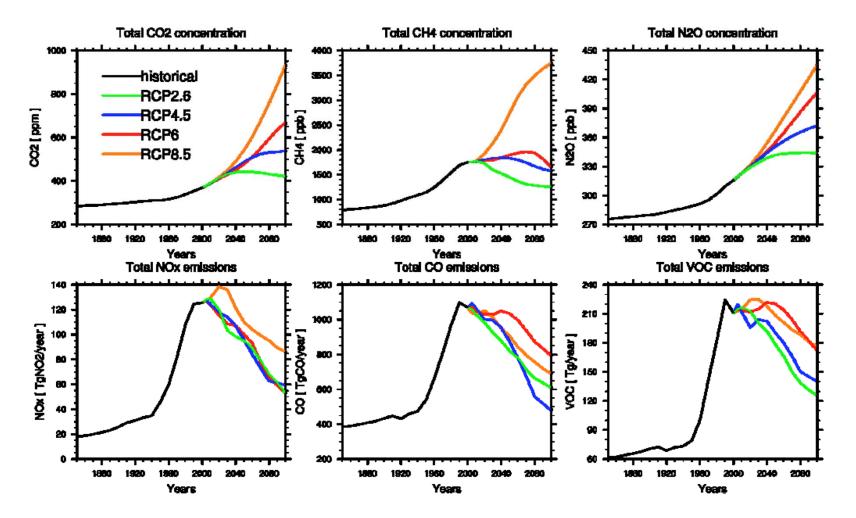
Surface ozone distribution (ppbv) calculated by LMDz-INCA for January and July present-day conditions.

Evolution of background tropospheric ozone since the pre-industrial.

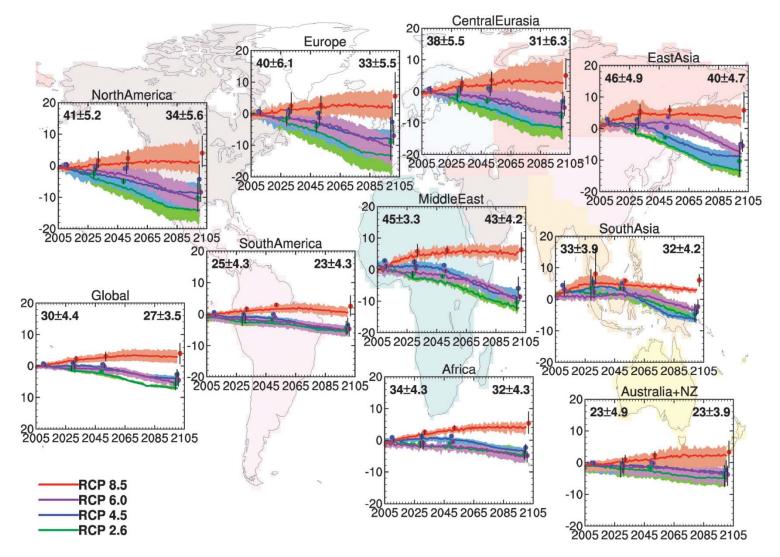




Evolution of tropospheric ozone in the free-troposphere



Evolution of future emissions of greenhouse gases and ozone precursors under the four *Representative Concentration Pathway* (RCP) scenarios from GIEC/AR5.

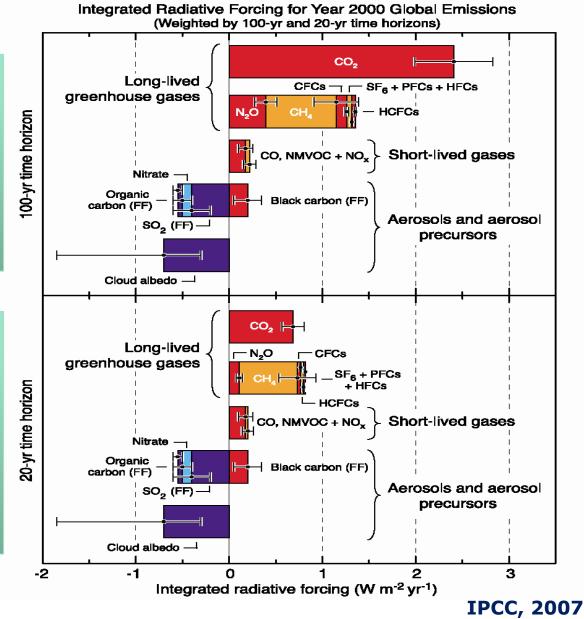


Global and regional future surface ozone levels (ppbv) simulated by the ACCMIP models for the four RCP scenarios.

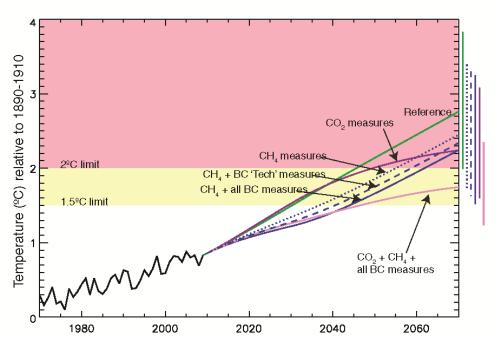
Fiore et al., 2012 Sino-French Institute for Earth System Science – PKU 8-12 April 2013 – 28

A 100-yr time horizon emphasizes the importance of longlived greenhouse gases (CO_2, N_2O, CH_4) emission reduction as far as long-term climate change is concerned.

In the meantime, a shorter time horizon (20 years) shows the importance of reactive species (ozone precursors) and aerosol emission control for a rapid benefic for climate.



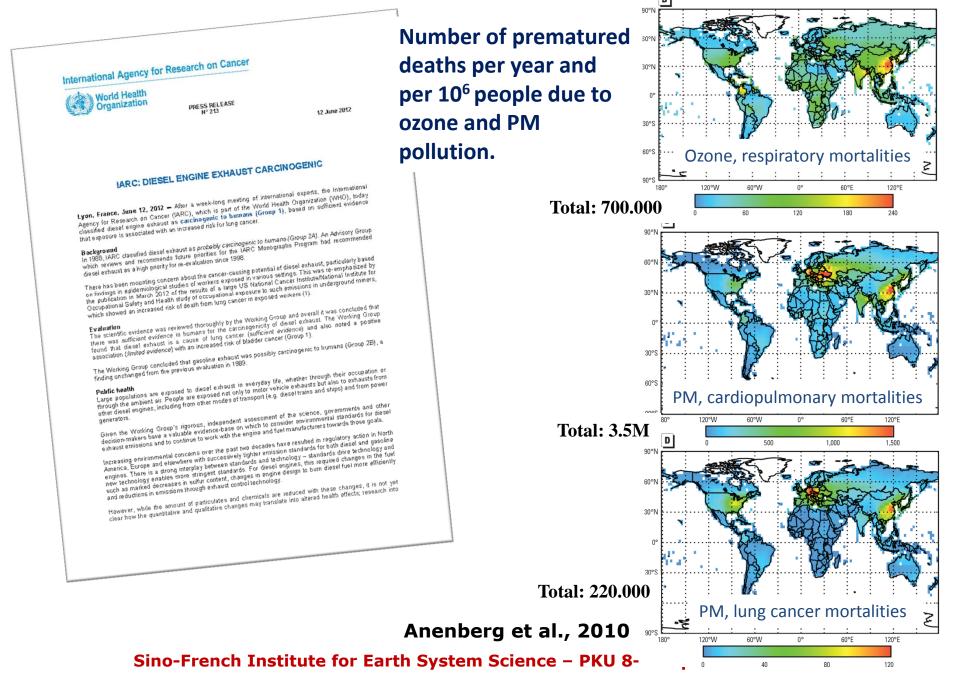
- More than 400 measures for reactive species (ozone precursors) and aerosol emission control have been considered.
- 14 key measures were selected as they improve both air quality and climate change. They involve CH₄ (and hence O₃) and BC.
- Measures acting only on CO₂ emissions cannot prevent the global warming to reach more than 2 degrees during the XXIst century because of the long residence time of carbon dioxide in the atmosphere.
- Action on both CO₂ and on short term climate agents such as CH₄ (and hence O₃) and BC, help to achieve this objective.



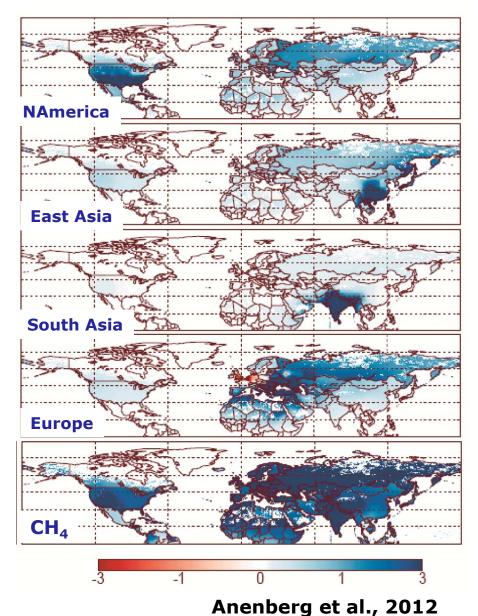
ig. 1. Observed temperatures (*42*) through 2009 and projected temperatures thereafter under various cenarios, all relative to the 1890–1910 mean. Results for future scenarios are the central values from nalytic equations estimating the response to forcings calculated from composition-climate modeling nd literature assessments (7). The rightmost bars give 2070 ranges, including uncertainty in radiative prcing and climate sensitivity. A portion of the uncertainty is systematic, so that overlapping ranges do ot mean there is no significant difference (for example, if climate sensitivity is large, it is large egardless of the scenario, so all temperatures would be toward the high end of their ranges; see www. iss.nasa.gov/staff/dshindell/Sci2012).

Shindell et al., 2012

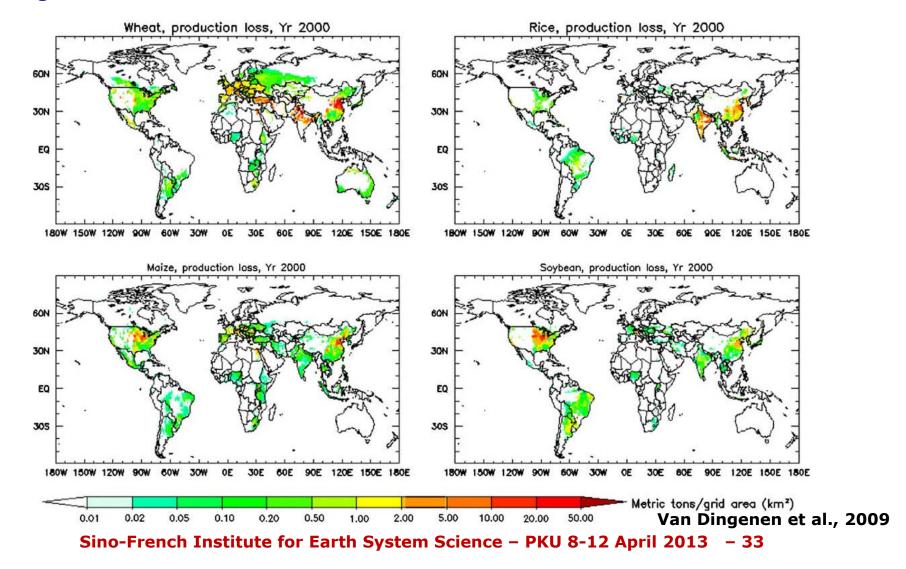
IPCC, 2007



- Annual avoided O₃ cardiopulmonary mortalities per million people resulting from 20% NO_x, NMVOC, and CO emission reductions in the region shown and a 20% global CH₄ mixing ratio reduction.
- Avoided mortalities in northern hemisphere when 20% reduction is applied in the different regions (hundreds). NH: 218; NA: 36; EU: 38; SA: 85; EA: 59.
- Domestic versus NH mortalities (hundreds). NA: 9 vs 36. EU: 17 vs 38; SA: 76 vs 85; EA: 43 vs 59.
- Methane reduction. More uniform reduction in mortality. NH: 160; NA: 11; EU: 39; SA: 48; EA: 38.



Present day ozone impact on four types of crops. Wheat: loss of 45 to 82 Mtons per year (7-12% of production). Rice: 17-23 Mtons (3-4%). Maize: 14-25 Mtons (3-5%). Soybean: 9-30 Mtonss (6-16%). Economic loss estimated globally 14-26 billion \$ per year. 40% of the damage in China and India.



Summary. Tropospheric ozone O_3 :

- 1. Is photochemically produced in the troposphere by the oxidation of methane CH₄, non-methane hydrocarbons NMHC, and carbon monoxide CO in the presence of nitrogen oxides NO_x.
- 2. Is a greenhouse gas contributing for about 0.35 W/m² to the present-day anthropogenic radiative forcing of climate.
- 3. Controls the oxidizing capacity of the atmosphere (OH) and hence affects the radiative forcing of other long-lived greenhouse gases such as CH_4 or aerosols such as sulfates.
- 4. Causes deleterious impacts to human health, including cardiovascular and respiratory mortality.
- 5. Can damage crops, leading yield reduction and deteriorating crop quality.
- 6. Can damage other land ecosystems and hence affect the capacity of plants to take up carbon dioxide from the atmosphere.